

# An Introduction to Radiative Forcing in Climate Change

Tatsuya Seiki @ Tomita team

# Self-introduction

断絶の2000年

	Before 2000	After 2000
部活	無し	ワンダーフォーゲル部
趣味	漫画	漫画、登山、沢登り、ボルダリング

1981年：北海道札幌市にて誕生

2000年：北海道から出たくて上京＋東京工業大学入学

2004年：雲が好きだったので修士から衛星観測の研究室@東京大学に編入

2010年：雲の数値モデルで学位（理学）取得

## 登った百名山（35個）

利尻岳、大雪山、トムラウシ山、十勝岳、鳥海山、会津駒ヶ岳、那須岳、男体山、日光白根山、両神山、雲取山、甲武信ヶ岳、金峰山、瑞牆山、富士山、谷川岳、草津白根岳、四阿山、浅間山、巻機山、白馬岳、五竜岳、鹿島槍ヶ岳、水晶岳、鷲羽岳、槍ヶ岳、穂高岳、常念岳、蓼科山、赤岳、甲斐駒ヶ岳、鳳凰山、塩見岳、光岳、石鎚山

# Cloud in Climate Study



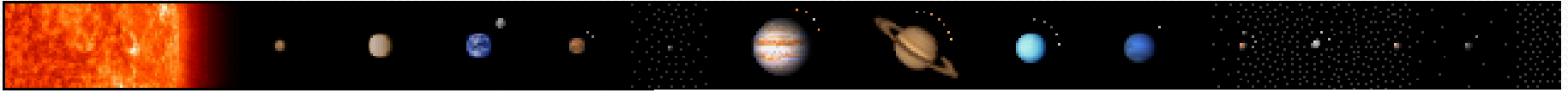
Various clouds = various brightness = different efficiencies of light-scattering (cooling the Earth)

	Thin	Mid	Thick
High	卷雲 Cirrus	卷層雲 Ciro-stratus	積乱雲 Cumulonimbus
Mid	高積雲 Alto-cumulus	高層雲 Alto-stratus	乱層雲 Nimbostratus
Low	積雲 Cumulus	層積雲 stratocumulus	層雲 stratus



# **Radiative Forcing and Climate Change**

# Radiative Forcing



## Radiative Forcing:

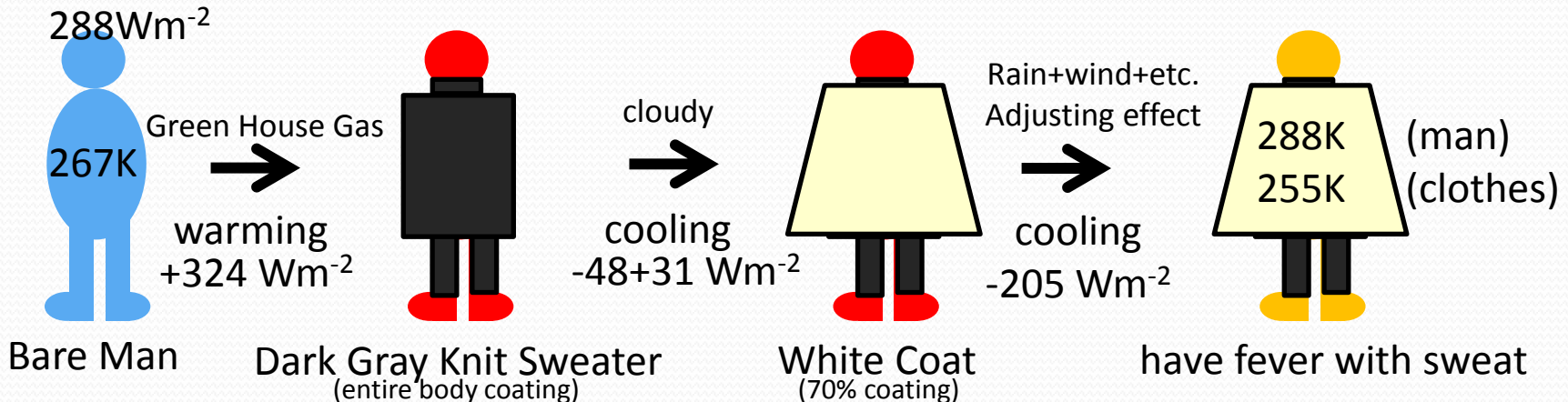
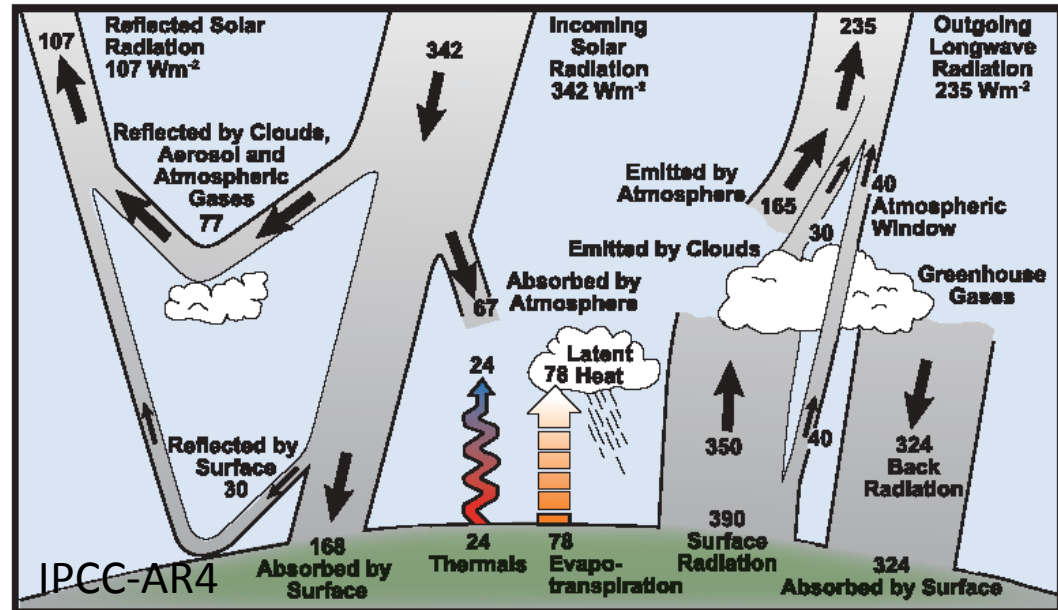
- Contribution to energy balance
- Concept to assess global warming

## Equilibrium condition

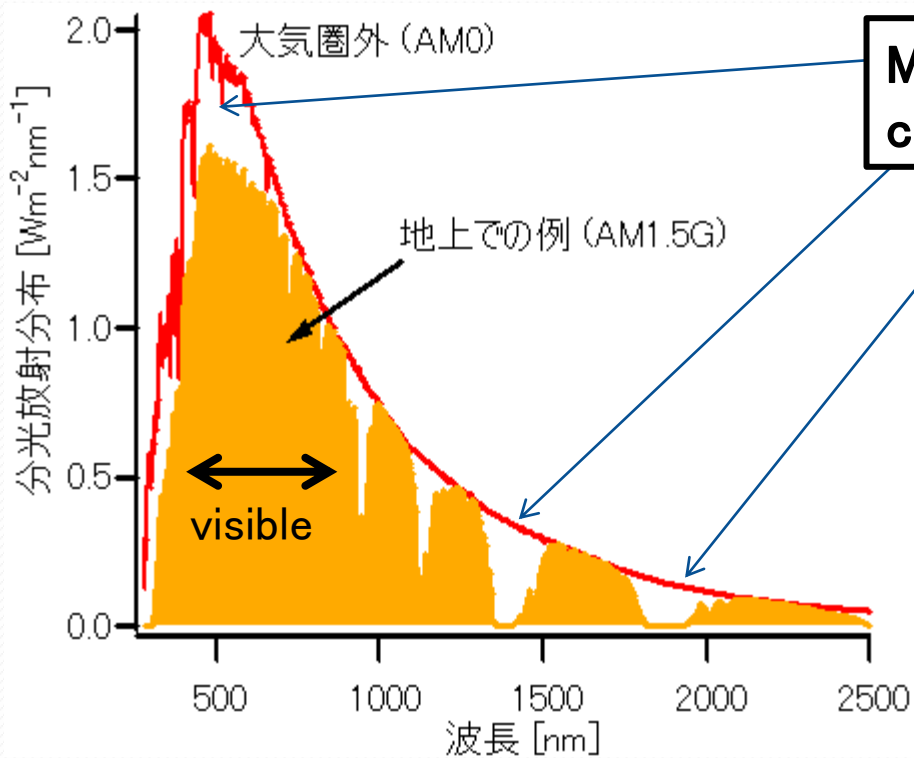
$$\pi r_E^2 (1 - \alpha) S_{sun} = 4\pi r_E^2 \sigma T^4$$

## Equilibrium + perturbation

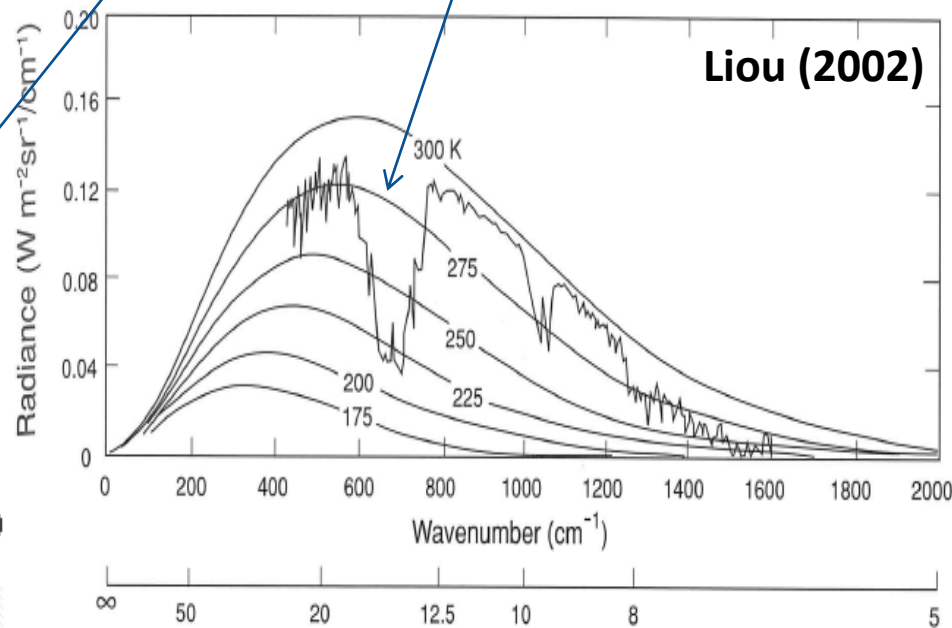
$$\pi r_E^2 (1 - \alpha) S_{sun} + \delta E = 4\pi r_E^2 \sigma (T + \delta T)^4$$



# Radiative Forcing (by wavelength)



**Missing energy comes from gas absorption  
c.f.) CO<sub>2</sub>, H<sub>2</sub>O, O<sub>3</sub>**

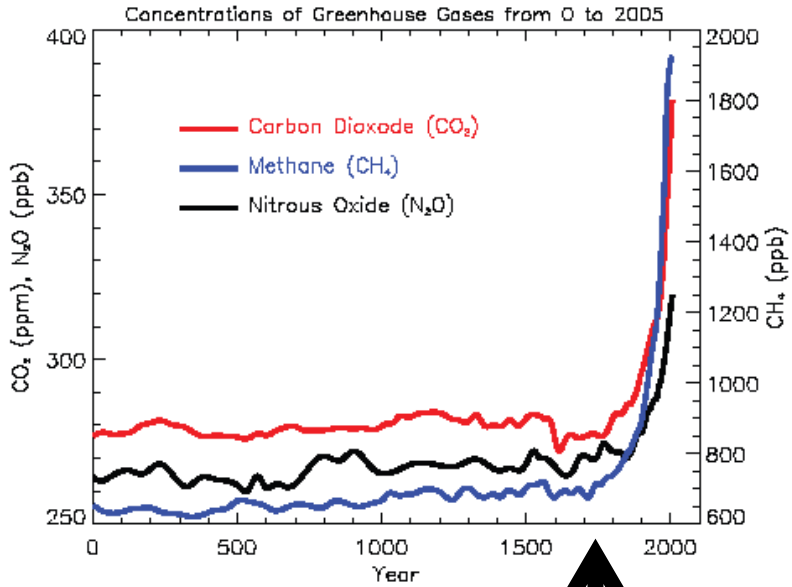


Radiative energy flux depends on wavelength,  
while optical properties of atmospheric gases/particles depends on wavelength.

$$B_{\lambda}(T) = \frac{2hc^2}{\lambda^5 \left( e^{hc/k_B\lambda T} - 1 \right)}$$

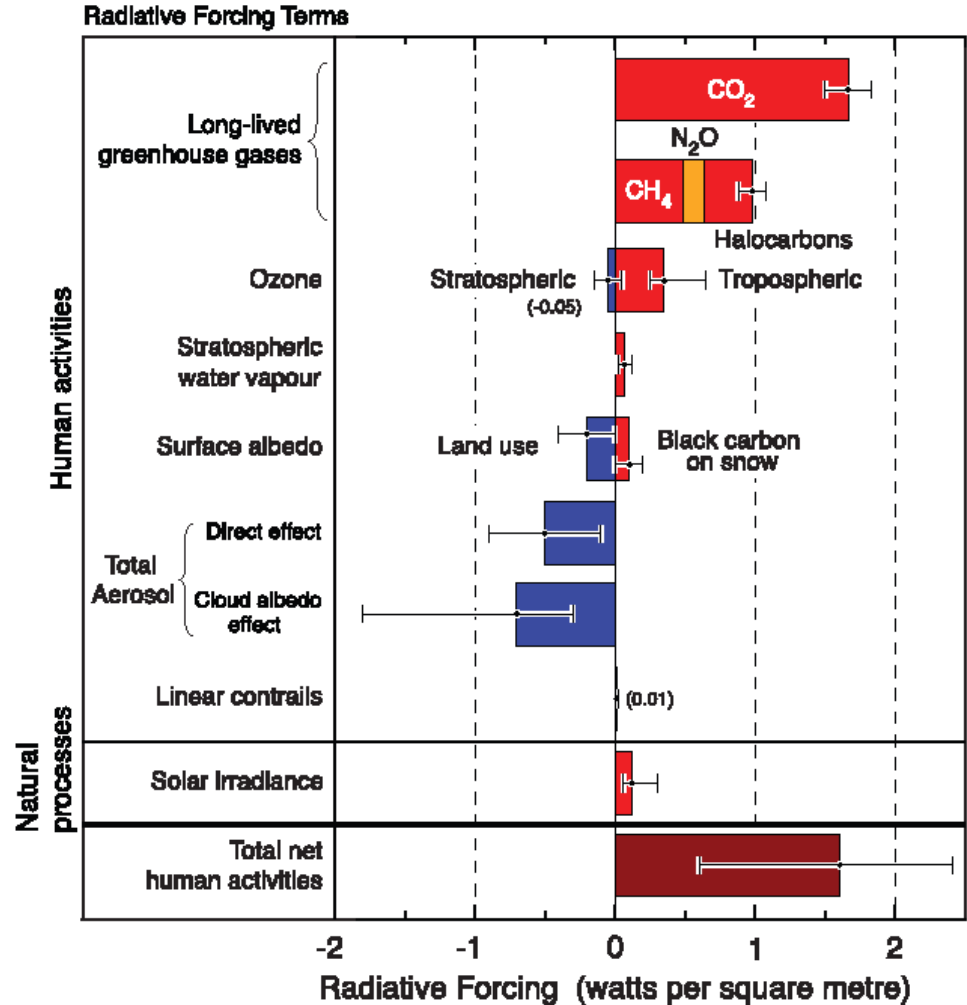
- ✓ Energy flux from Solar beam: H<sub>2</sub>O, O<sub>3</sub>, Cloud, Aerosol
- ✓ Energy flux from Earth heat: H<sub>2</sub>O, CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O, O<sub>3</sub>, Cloud

# Anthropogenic Radiative Forcing



Industrial Revolution: Steam Engine by Newcomen, 1712; Watt, 1769

Radiative forcing of climate between 1750 and 2005



Human activities role as an additional radiative forcing into climate systems after 1700s. Greenhouse gases have strong impact to heat the Earth with less uncertainties, while aerosols has strong impact to cool the Earth with much uncertainties.

# Aerosol Radiative Forcing

2007.6.17-21, Beijing (photo by Seiki)



At Beijing, it is likely to be foggy even though it is not cloudy. Aerosols (Atmospheric Pollution) overcast sky to decrease energy exp(-1) near the ground, particularly around industrial regions.

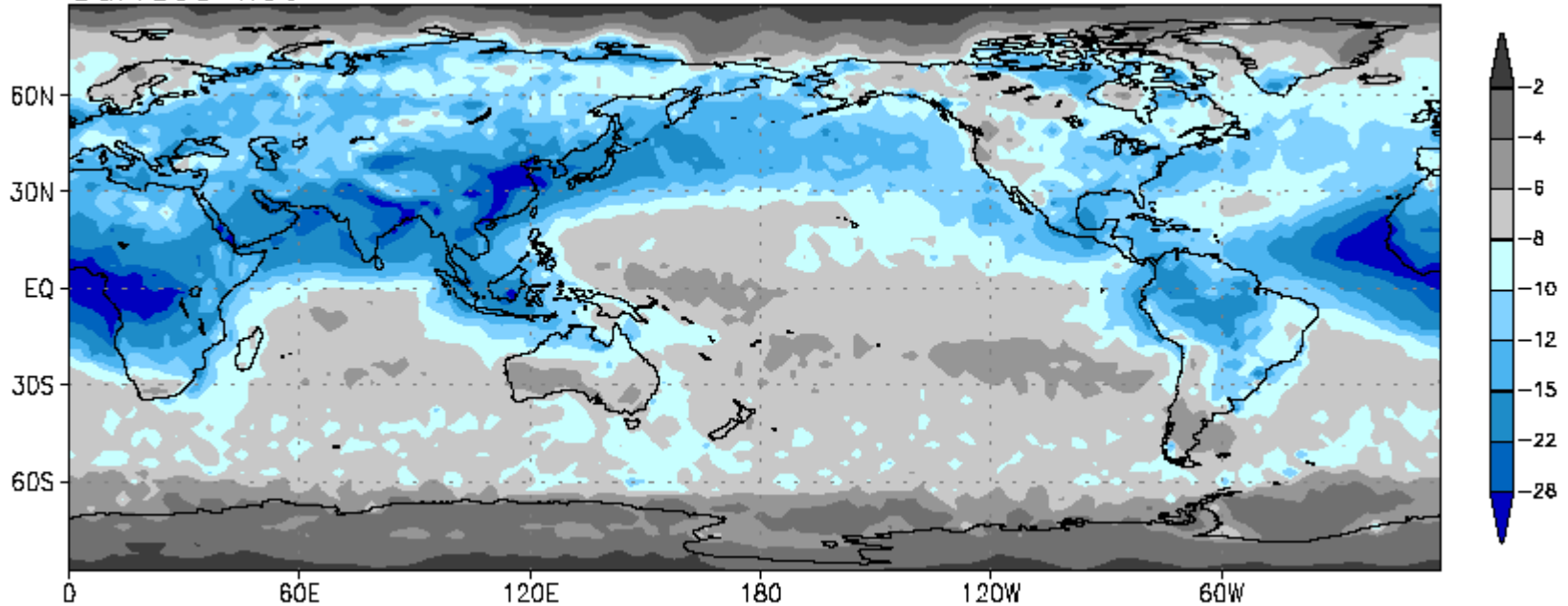


# Global dimming by Aerosols

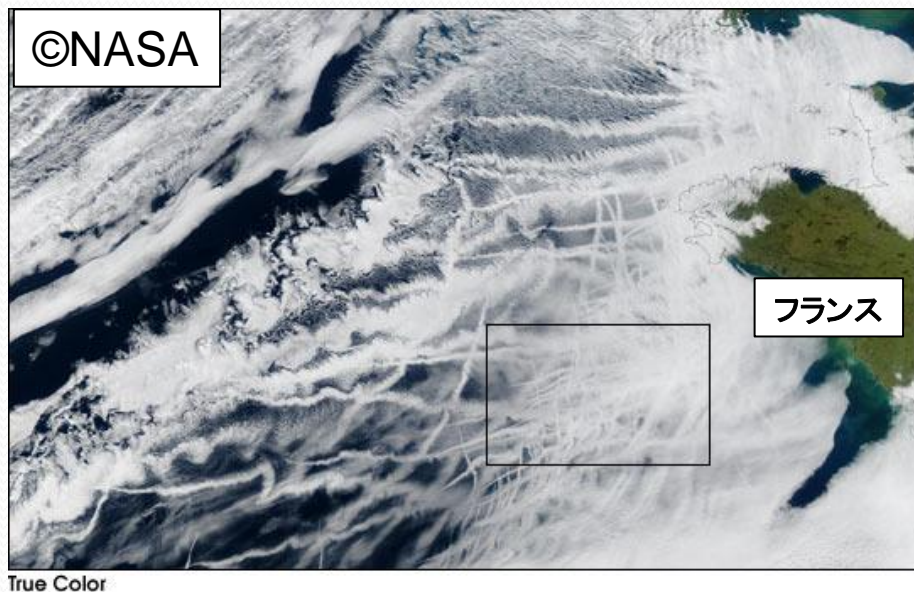
Aerosol radiative forcing at the surface by multi-satellite/ground observations.

**Kim and Ramanathan (2008)**

Surface net



# Aerosol Indirect Radiative Forcing



## 典型的な雲

$1\text{m}^{-2}$ あたりの質量～  $50\sim 100\text{ g m}^{-2}$

太陽光の散乱～雲の  $\text{断面積}(\pi r^2)$  に比例

50gの水滴一個だと半径は2.3cm  
断面積は $16.4\text{cm}^2$

→ $1\text{m}^2$ の地面のうち、0.164%しか覆われない

海上の雲粒の典型的な大きさ～  $20 \times 10^{-6}\text{ m}$

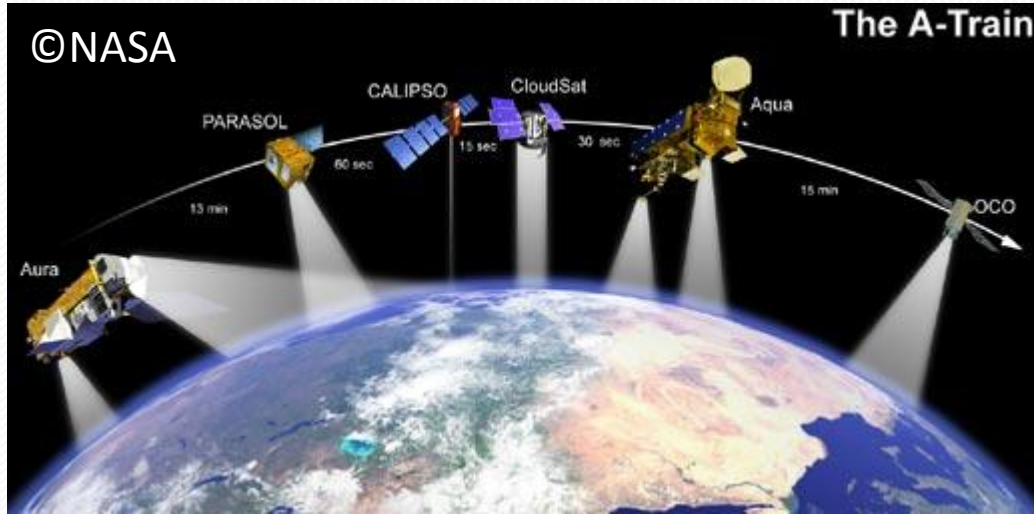
→粒子数は $1.5 \times 10^9$ 個、断面積は～  $1.88\text{ m}^2$

→188%も雲が覆う！

粒子サイズが1/2になると断面積は2倍になる！

人間の産業活動は硫酸などのエアロゾル(大気汚染物質)を大量に排出し、エアロゾルが水蒸気をまとめて多量の雲凝結核となる。  
エアロゾルが多い地域では雲粒が小さくなって太陽光を反射しやすくなっている

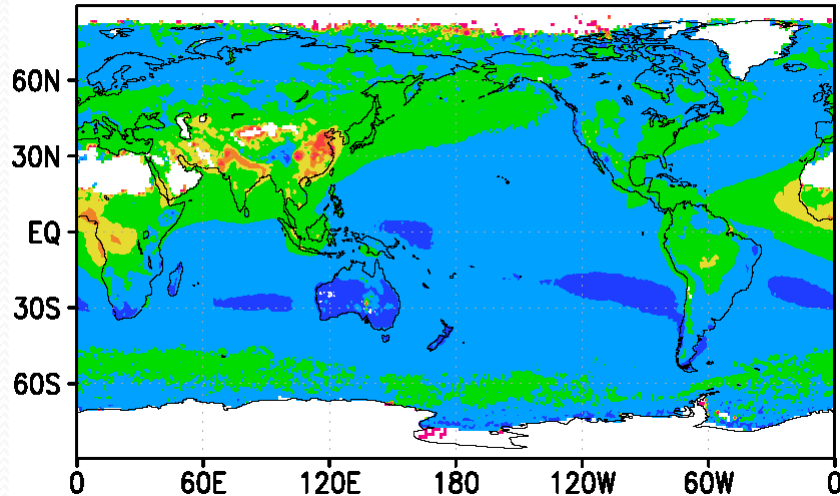
# Aerosols Distribution by Satellites



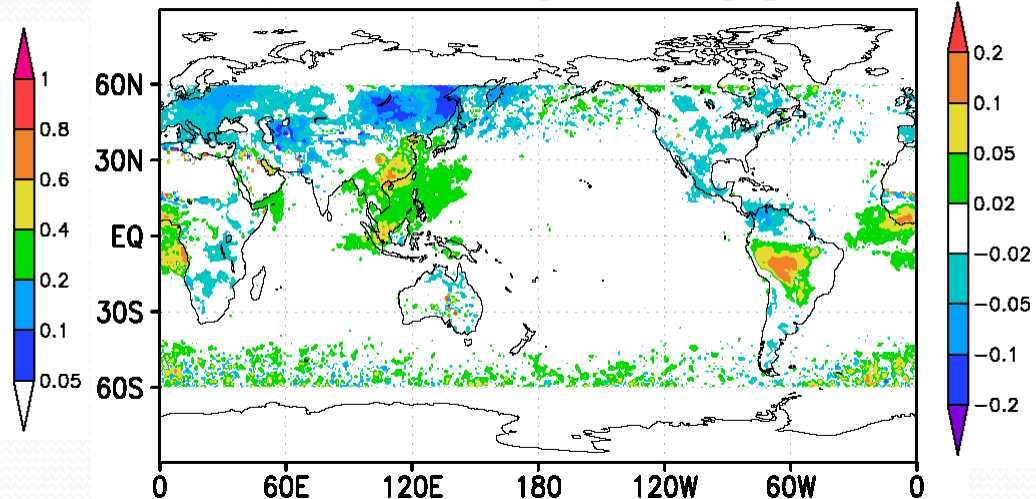
After 1980s, NASA has led observation projects using satellite remote sensing. (e.g., ISCCP, ERBE, A-Train)

We can estimate radiative forcing by Cloud and aerosols quantitatively.  
➔ Improving Climate Models

Aerosol Optical Thickness [2000:2008]

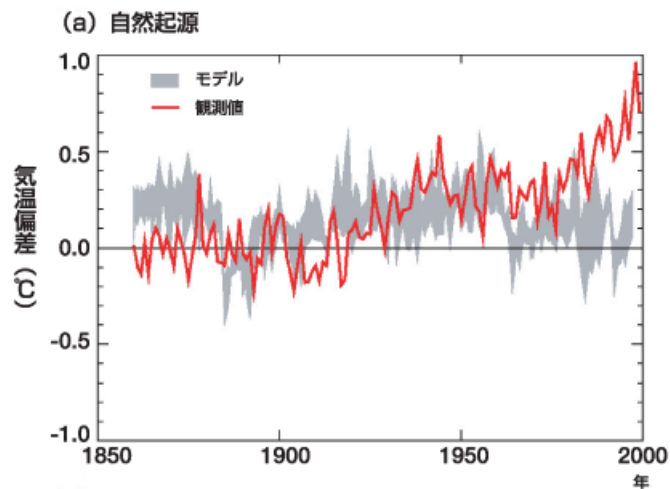


Aerosol Optical Thickness [2005:2008] - [2001:2004]

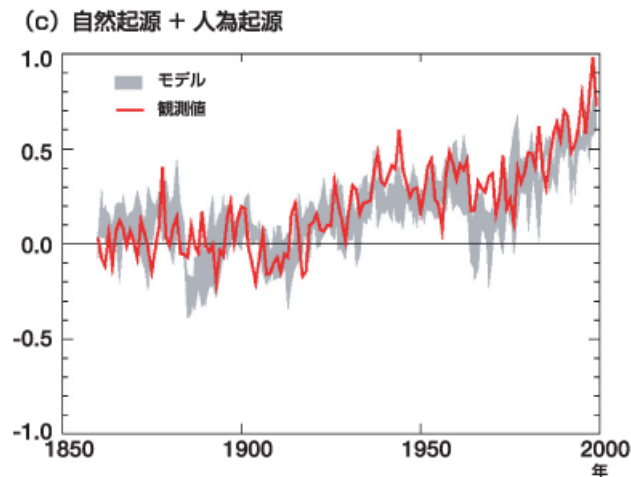


MODIS/Aqua-Terra

# Global Warming Projection



自然起源の応答（太陽放射や火山噴火）だけでは、20世紀後半の温暖化は説明できない。



すべての人為起源と自然起源の因子を複合させると、過去140年間の観測値とモデル計算が最もよく説明される。

## B1「持続的発展型社会シナリオ」

- ・環境の保全と、経済の発展を地球規模で両立する。

## A1T: 非化石エネルギー源を重視(新エネルギーの大幅な技術革新)

## B2「地域共存型社会シナリオ」

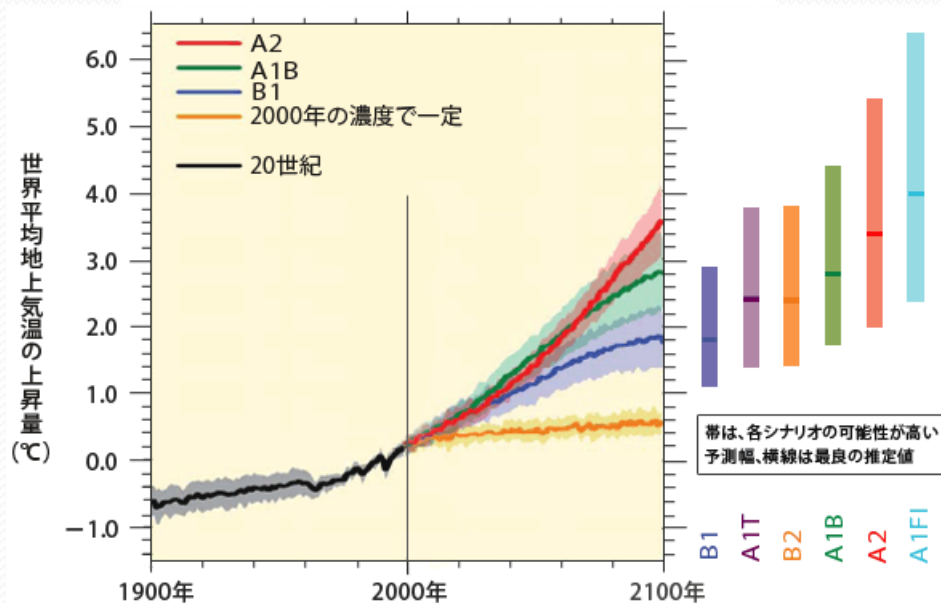
- ・地域的な問題解決や世界の公平性を重視し、経済成長はやや低い。
- ・環境問題等は、各地域で解決が図られる。

## A1B: 各エネルギー源のバランスを重視

## A2「多元化社会シナリオ」

- ・世界経済や政治がブロック化され、貿易や人・技術の移動が制限。
- ・経済成長は低く、環境への関心も相対的に低い。

## A1FI: 化石エネルギー源を重視





# **General Circulation Model**

# General Circulation Models (GCMs)

Observations (local/global, long/short, accurate/coarse)

empirical

Approximation, assumption

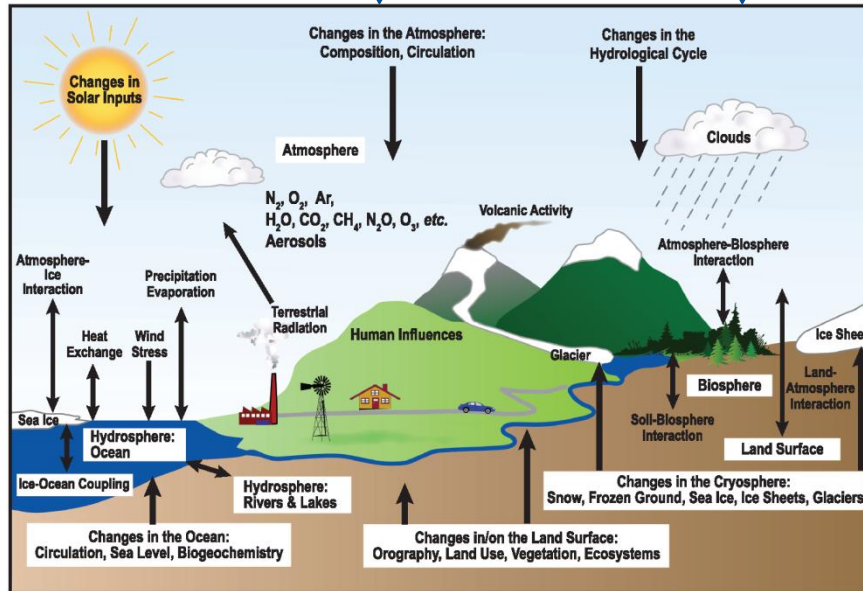
Approximation, assumption

Boundary Condition

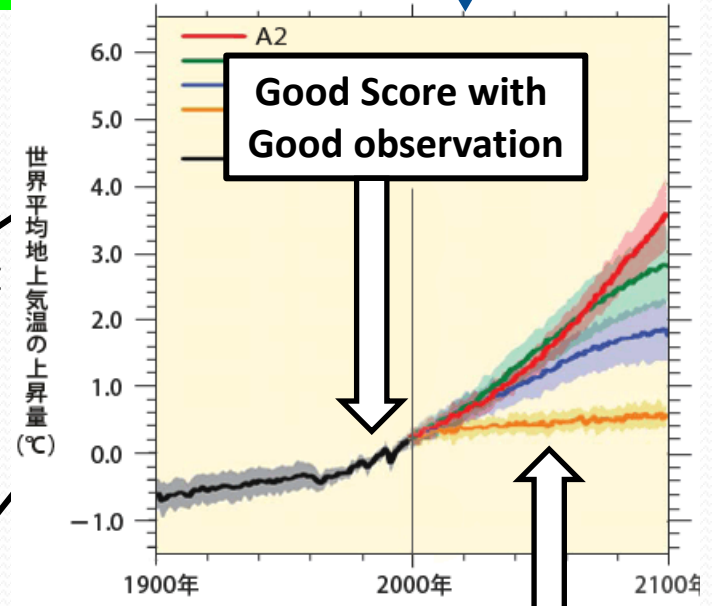
Initial Condition

Correction optimization

Comparison, Evaluation



predict



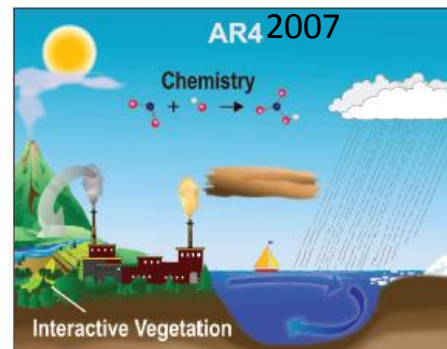
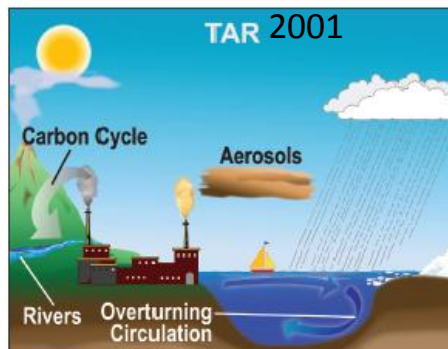
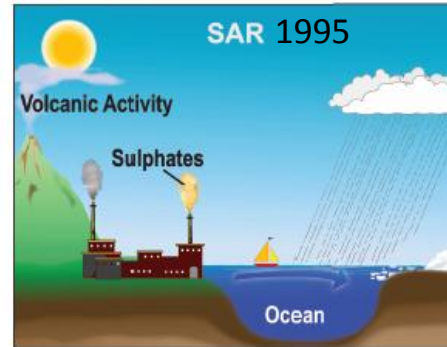
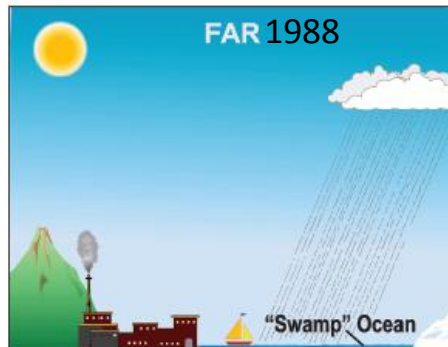
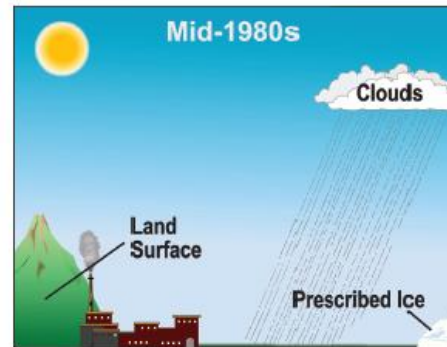
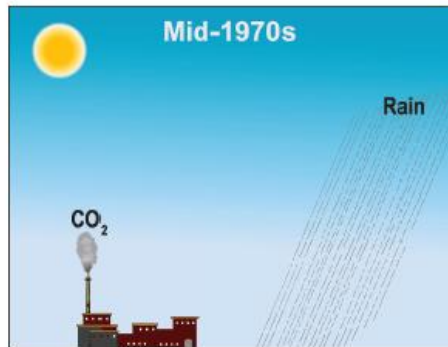
approximation

Adding component/interaction

Theory (Component1, 2, 3, 4, ...)

Poor score because of human activity with much uncertainties

# Development of GCMs



General Circulation Models contain:

➤ Geophysical Fluid Dynamics  
(fluid+thermo+gravity+Coriolis+salinity)

- Atmospheric circulation
- Oceanic circulation
- Phase change/Latent heat

• Vapor ↔ cloud,rain

- Sea-water ↔ sea-ice, salinity

➤ Land Surface

- Vegetation, urbanization, cultivation
- River
- Soil (heat/water transport)
- Ice-sheet

➤ Radiation (e.g., Solar Incidence)

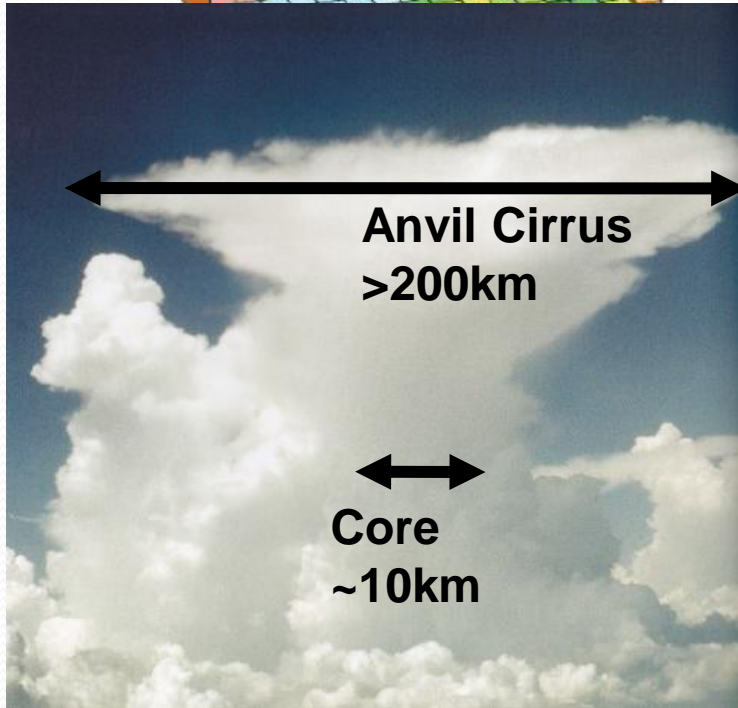
- Gas absorption
- Rayleigh Scattering
- Mie Scattering by Cloud, Aerosol

➤ Chemistry (e.g., ozone hole)

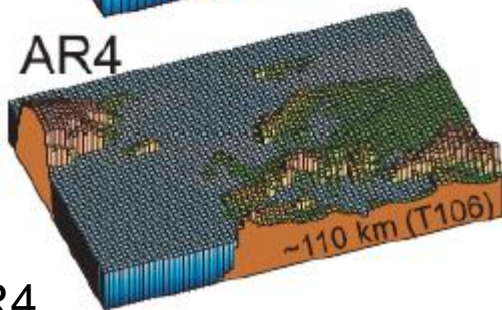
- Chemical Reaction
- Wet deposition
- Emission

# Advances in Simulation

1988 FAR

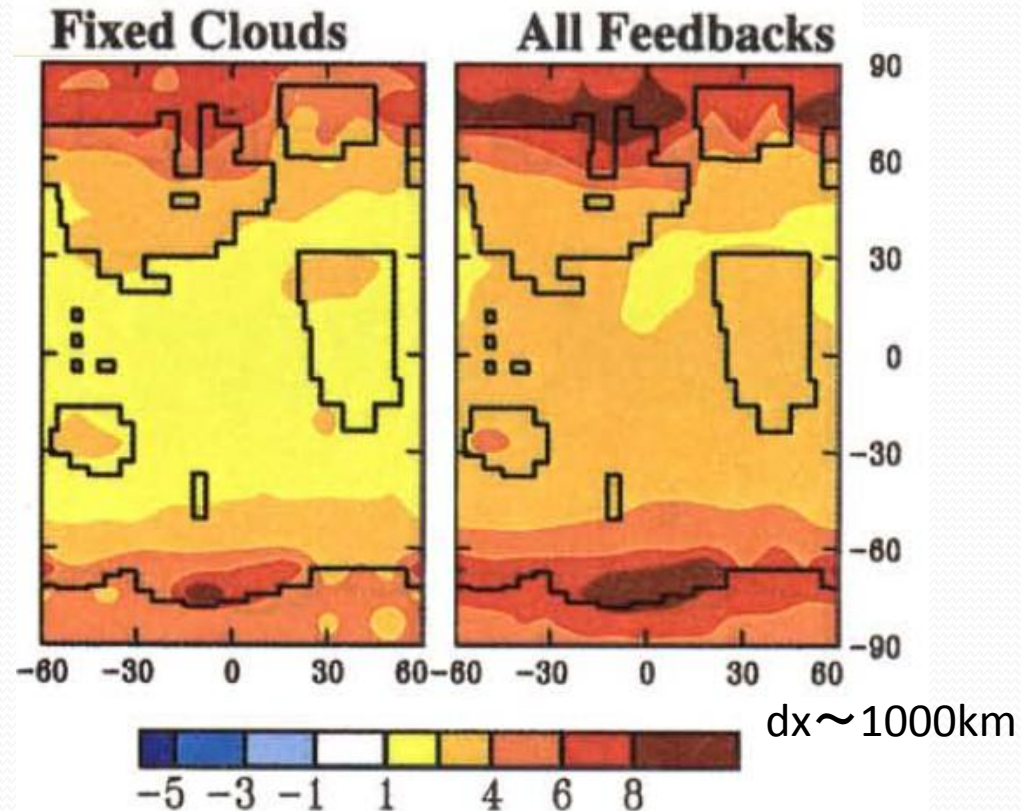



2007 AR4



IPCC-AR4

$2 \times \text{CO}_2 \Delta T$  ( $^{\circ}\text{C}$ )



Hansen et al. (1997)

1988, Dr. James Hansen@NASA suggest Global Warming

1988, IPCC is established

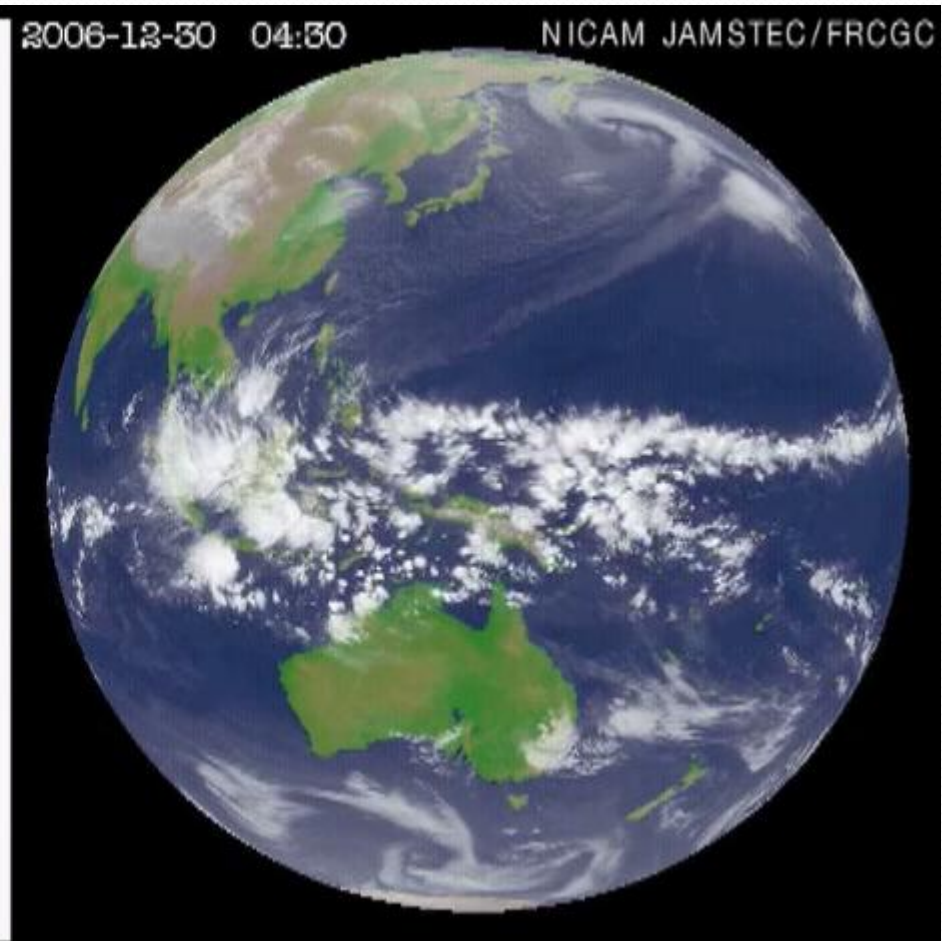
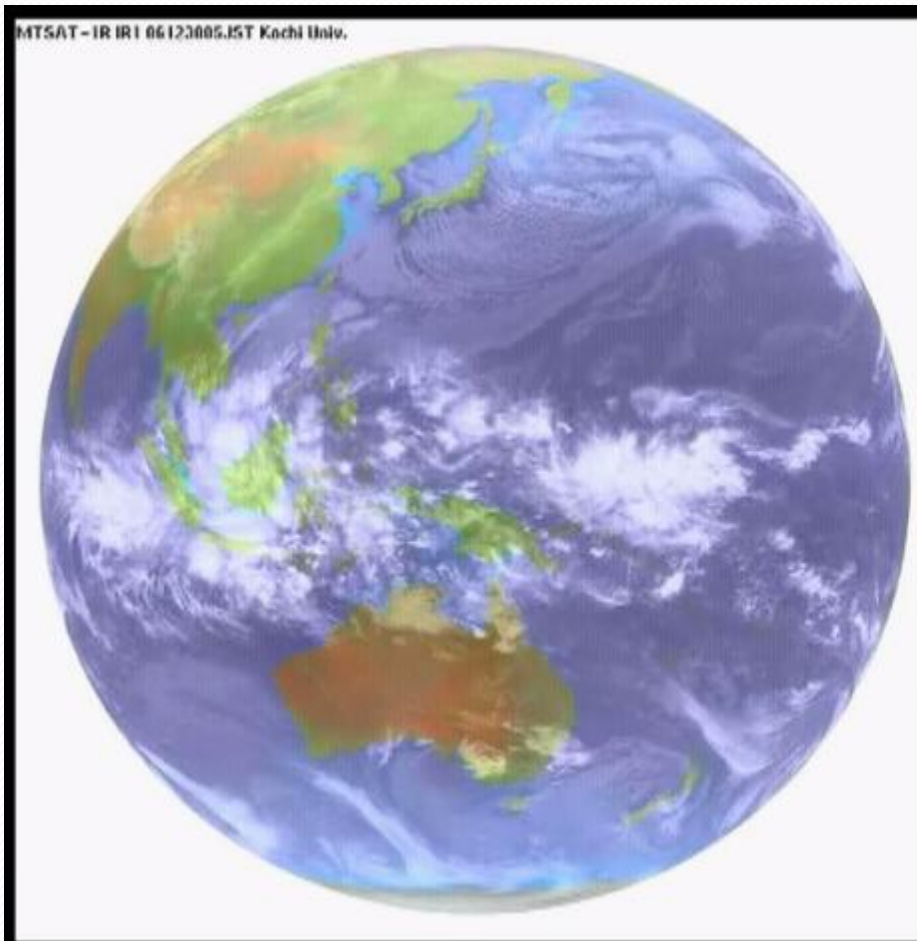
Recently, it is found that poorly resolved cloud causes the most severe error in climate change.



# Global Cloud Resolution Model, NICAM @ The Earth Simulator, Japan

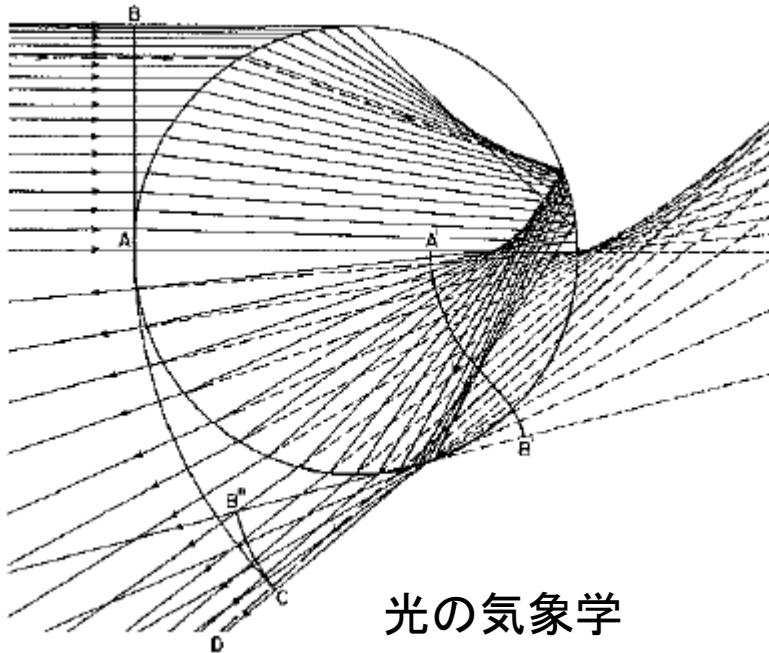
Reproduction of Madden Jullian Oscillation (Miura et al., 2007)

→ horizontal resolution=3.5km (↔ 30000 times heavier than conventional ones: dx~110km)



# Cloud Optical Properties (classic)

Descartes Rainbow is introduced in “Discourse on Method (方法序説)”(1637)



Snell's law: refraction theory

$$\frac{\sin \theta_A}{\sin \theta_B} = \frac{v_A}{v_B} = \frac{m_A}{m_B} \quad m: \text{refractive index}$$

Application of Maxwell's law to sphere

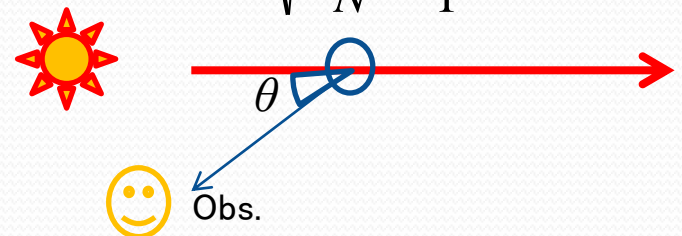
$$\begin{aligned} E_l^i &= R_l E_l^i, & E_l^t &= T_l E_l^i, \\ E_r^i &= R_r E_r^i, & E_r^t &= T_r E_r^i. \end{aligned}$$

*s.t.*

$$\begin{aligned} R_r &= \frac{\cos \theta_i - m \cos \theta_t}{\cos \theta_i + m \cos \theta_t}, & R_l &= \frac{m \cos \theta_i - \cos \theta_t}{m \cos \theta_i + \cos \theta_t}, \\ T_r &= \frac{2 \cos \theta_i}{\cos \theta_i + m \cos \theta_t}, & T_l &= \frac{2 \cos \theta_i}{m \cos \theta_i + \cos \theta_t}. \end{aligned}$$

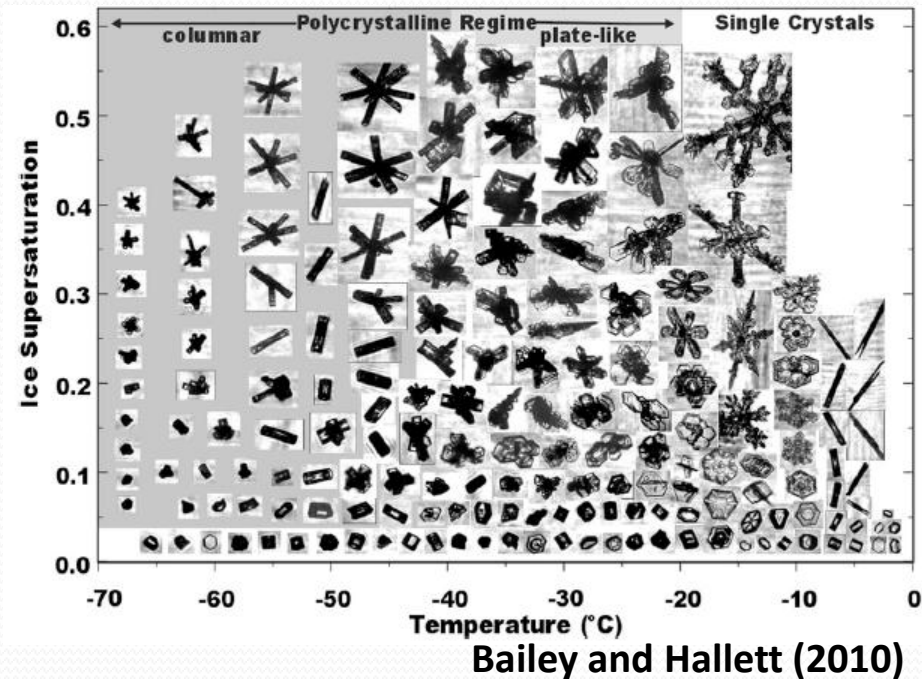
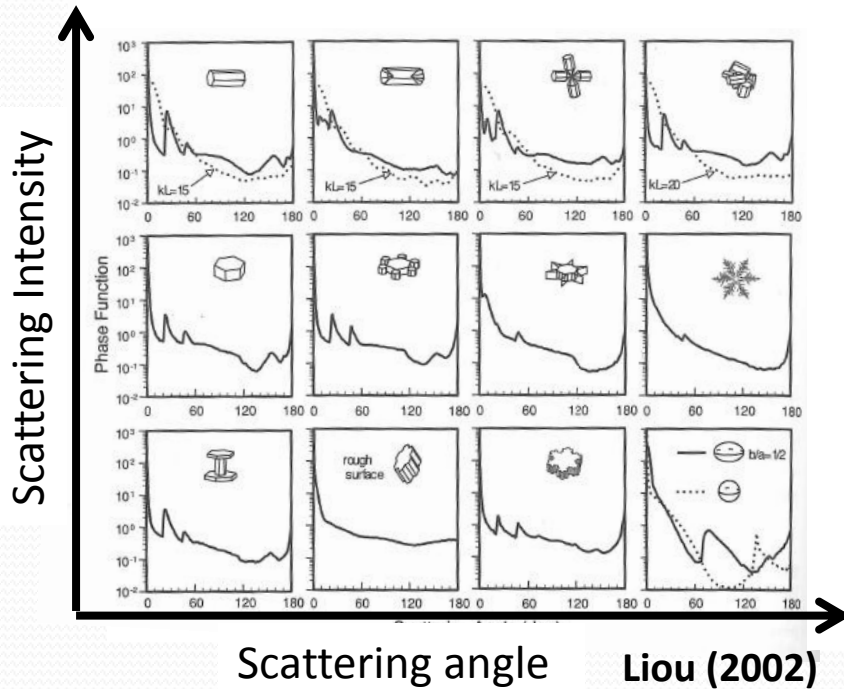
Deriving a major angle to concentrate energy fluxes after Nth refractions.

$$\theta_i = \cos^{-1} \sqrt{\frac{m_{\text{water}}^2 - 1}{N^2 - 1}}$$



color	purple	yellow	red
$m_{\text{water}}$	1.3435	1.3341	1.3318
$\theta_i$ (deg)	58.80	59.35	59.48
$\theta$ (deg)	40.58	41.91	42.25

# Cloud Optical Properties (advanced)



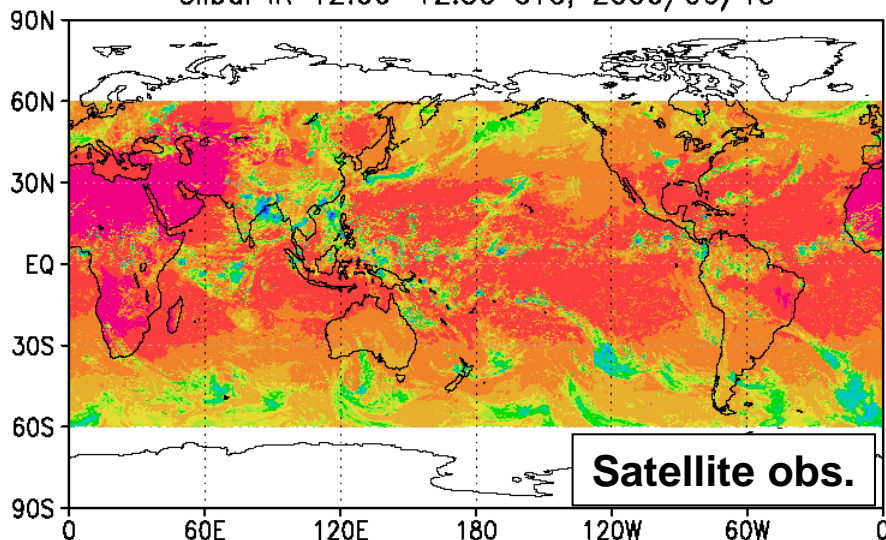
## Current topics of my research

1. Ice Nucleation Modeling:  
Aerosol/Supercooled liquid water → Ice Crystals
2. Estimation of the Impact of Non-sphericity of ice crystals

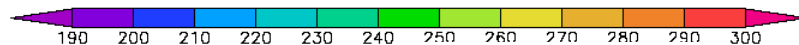
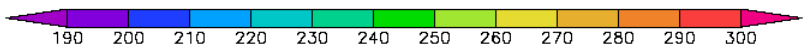
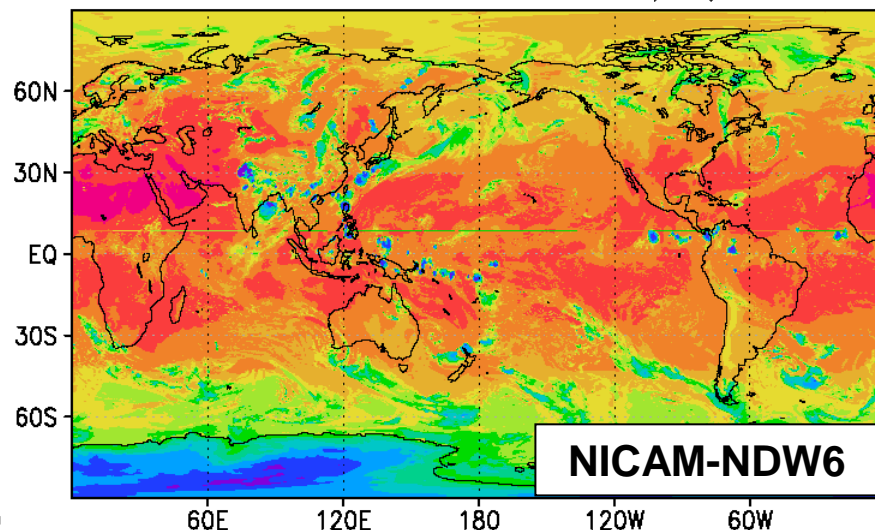
# Global Simulation (now undergoing)

## Infrared Brightness Temperature [K]

Global IR 12:00–12:30 UTC, 2006/06/15

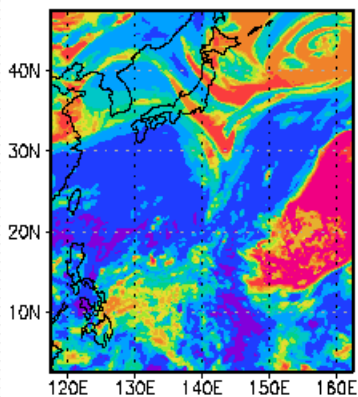


NICAM-NDW6, 12:00 UTC, 2006/06/15

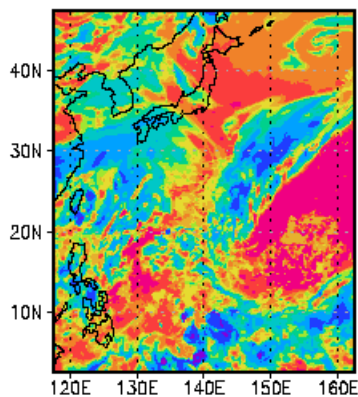


Differences by schemes

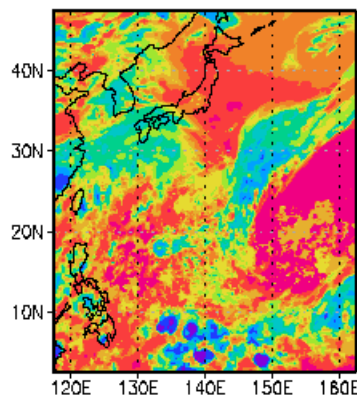
Meyers+Hex



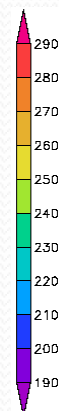
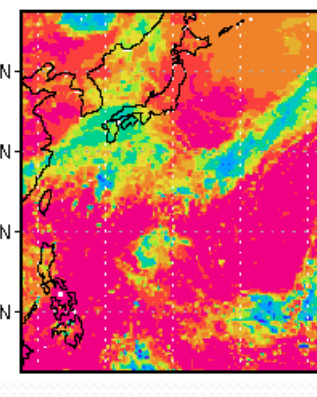
Phillips+Hex



Phillips+Col



Satellite obs.



Ice Nucleation Scheme: Meyers et al. (1992) or Phillips et al. (2006)  
Ice Crystal Shape: Hexagonal Plate (Hex) or Hexagonal Column (Col)

# Climate Simulation in the Future (unlimited...)

None knows the sufficient spatial resolution,  
and the sufficient combination of cloud/aerosol interaction.  
(Do uncertainties converge or not in the future ?)

- GCMs need fine spatial resolution to resolve cloud as much as possible.
  - $dx \sim 3.5\text{km}$  is the finest at this moment.
  - ➔  $dx > 1\text{km}$  is challenging in the next step  
(more than  $3.5^3$  times heavier)
- It is better to treat many species of cloud/aerosol particle as much as possible.
  - ➔  $\text{Cost} \sim N^2 \times M^2$ :  $N/M$  is the number of cloud/aerosol.  
( $N=5$ ,  $M=24$  at this moment)