Using sequential DA to identify and fix surface flux bias

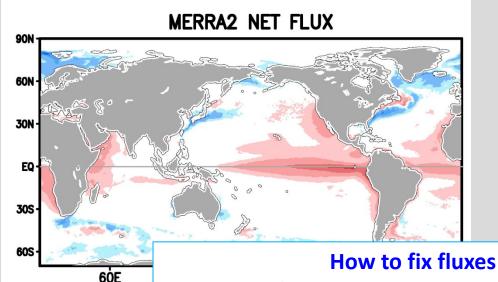
- Current approach to flux estimation the problem
- Temperature increments from SODA ocean reanalysis
- Fluxes that reduce increments + a fitness test
- Different bulk parameterizations and different atmospheric reanalyses affect the results

Jim Carton, G. Chepurin, L. Chen, and S. Grodsky
Many thanks to fellow incrementalists: Eugenia
Kalnay and Kriti Bhargava

Reanalysis fluxes and their differences



- MERRA2
- ERA-I
- JRA55



Meteorological approach:

Fix each component *a posteriori*:

Large and Yeager (2009); Brodeau et al. (2010)

Parameter estimation:

Kang and Kalnay (2013); Bateni and Entekhabi (2012)

Oceanographic approach:

Profiles as a constraint:

- 0) DaSilva (1994); Lamb (1981); Hastenrath & Merle (1983)
- 1) Isemer et al. (1989); Grist & Josey (2003)
- 2) Stammer et al. (2004)

SODA3

$$h\rho C_p \frac{D\theta}{Dt} \cong Q$$

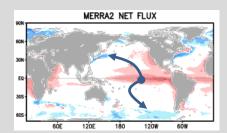
Vertically averaged energy budget





$$h\rho C_{p}\nabla\cdot\overline{\mathbf{u}\,\theta}\cong Q$$

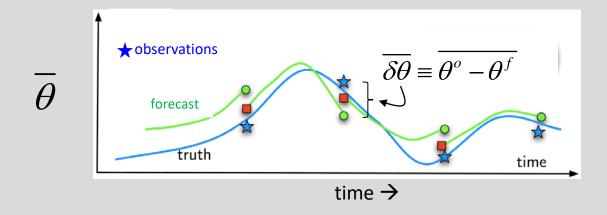
$$h\rho C_p \partial \overline{\theta} / \partial t \cong Q$$



$$h\rho C_p \partial \overline{\theta} / \partial t + L\rho \partial h_i / \partial t \cong Q$$
 If you account for sea ice

Sequential data assimilation

Example: vertically averaged temperature

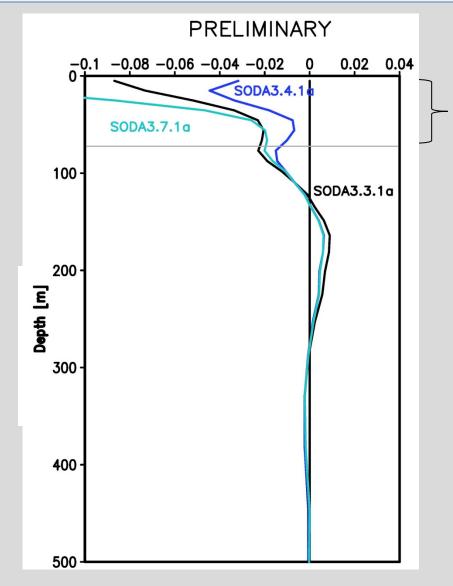


Incremental internal energy equation (averaging over an assimilation cycle)

$$\widehat{\overline{\delta\theta}} + \widehat{\overline{\delta\mathbf{u}} \cdot \nabla \theta^f} + \widehat{\mathbf{u}^f \cdot \nabla \delta\theta} \cong [\widehat{Q}^o - \widehat{Q}^f] \frac{\Delta t}{\rho C_p h}$$

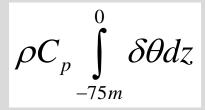
$\widehat{\overline{\mathcal{S}\theta}}$ averaged 60°S-60°N, 2007-2014

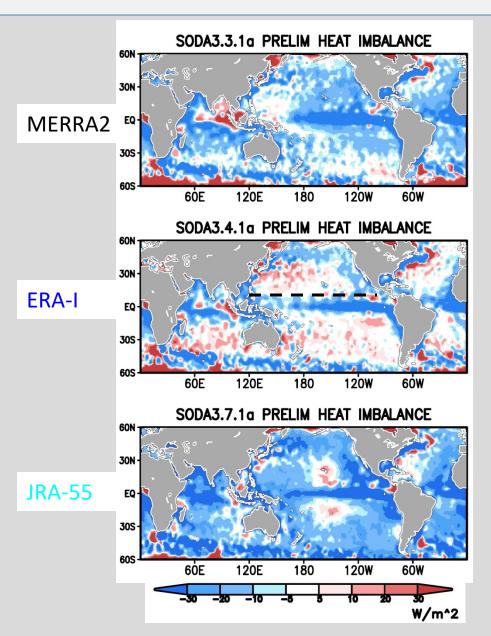
- MERRA2
- ERA-I
- JRA-55



Most of the misfit is in the upper 75m

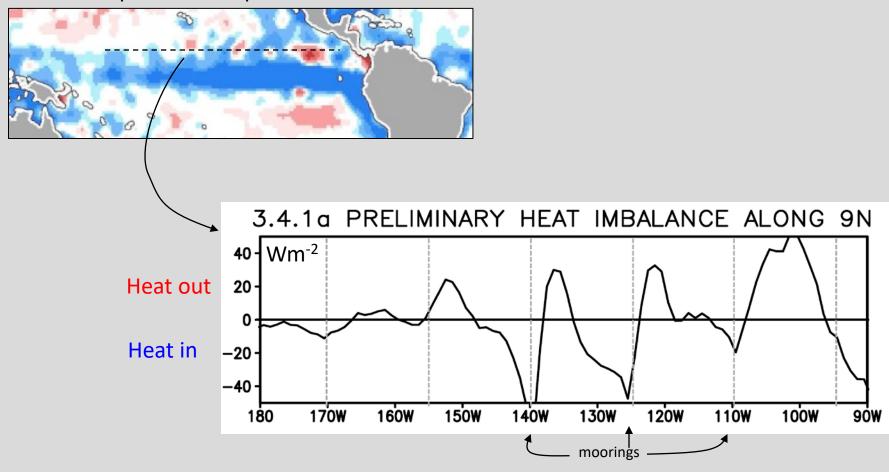
Average heat imbalance (2007-2014)

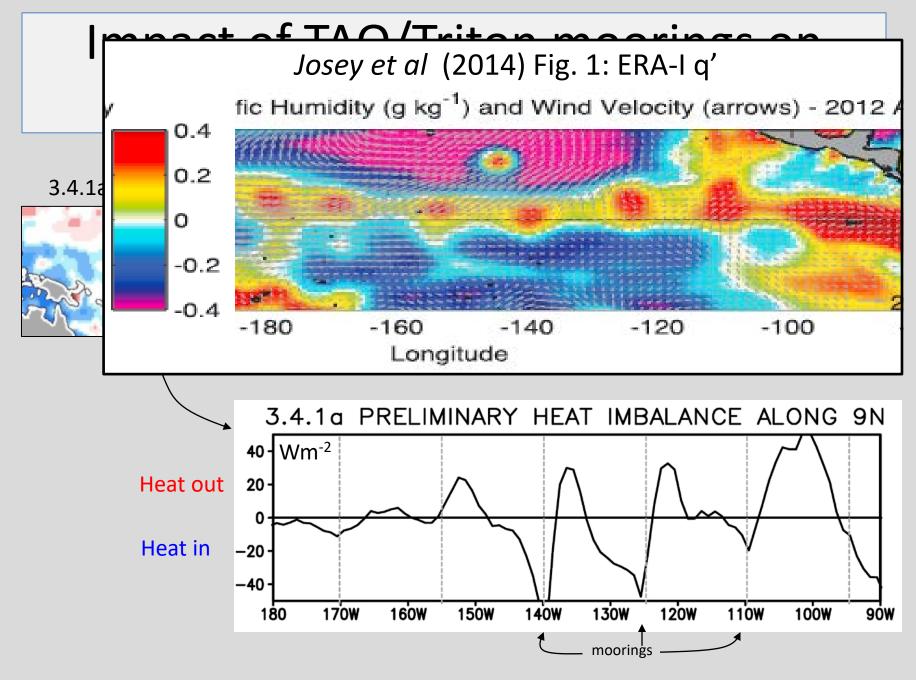




Impact of TAO/Triton moorings on ERA-I fluxes

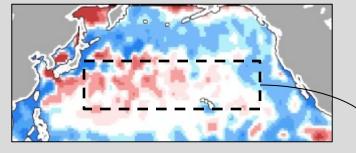
3.4.1a Equatorial temperature imbalance





Impact of seasonal cycle

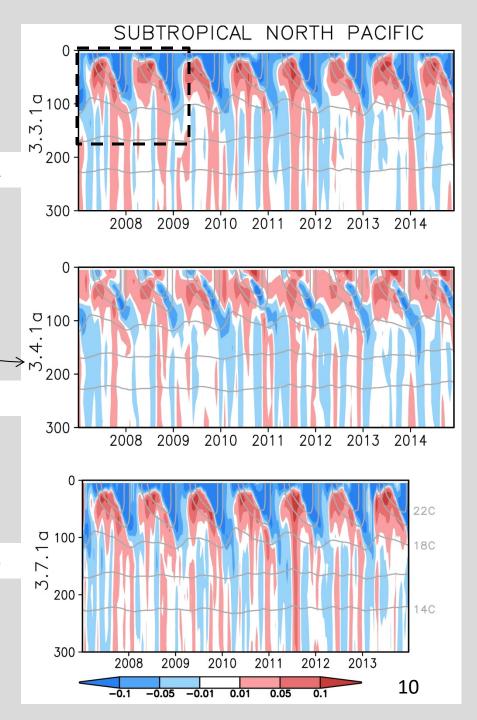
3.4.1a NP temperature imbalance



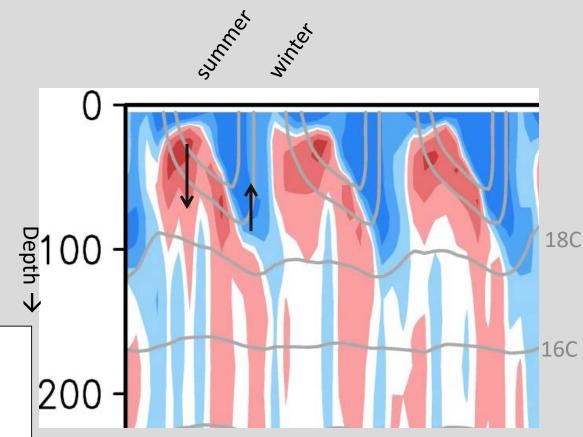
MERRA2

ERA-I

JRA-55



Detail of the 3.3.1a North Pacific temperature imbalance



What impact on mixed layer?

- Deepens summer thermocline,
- Weakens ΔT at the base of the winter thermocline

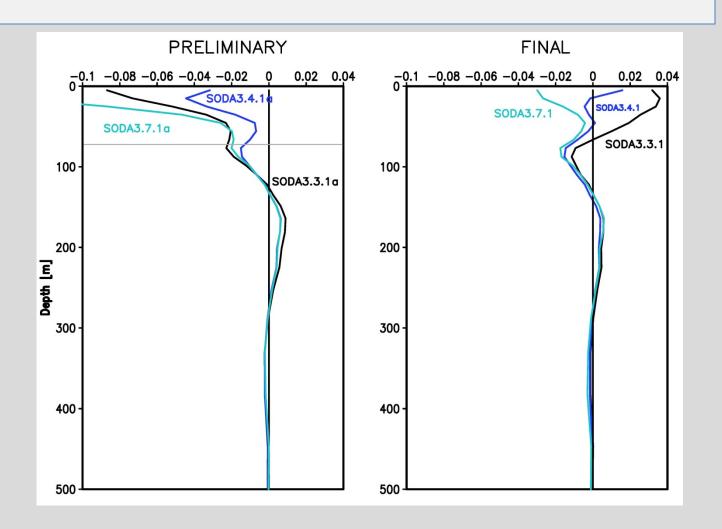
Modify net surface heat flux in the SODA forecast model

$$\frac{D\mathbf{u}}{Dt} = \dots$$

$$\frac{DS}{Dt} = \dots$$

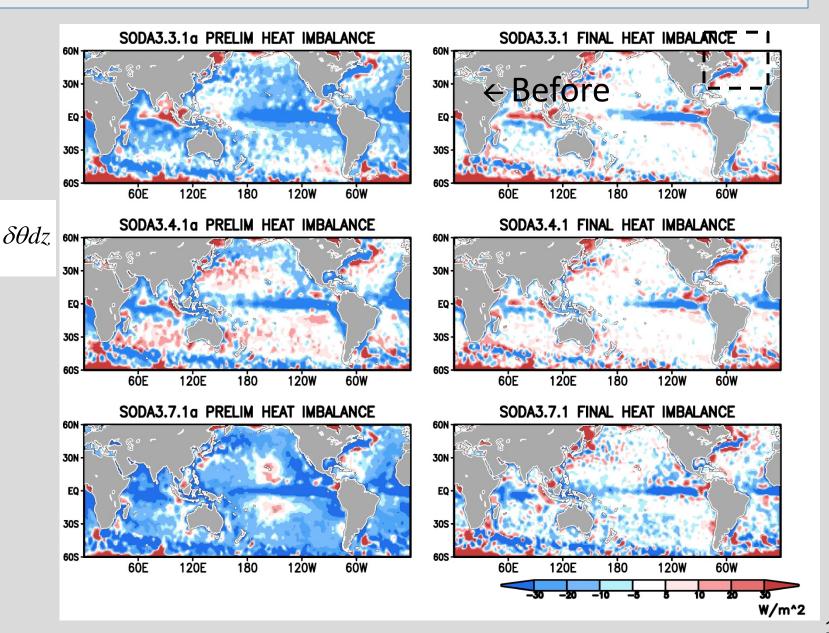
$$\frac{D\theta}{Dt} \cong d(z) \frac{Q + \delta Q}{\rho C_p}$$

$\widehat{\overline{\mathcal{S}\theta}}$ averaged 60S-60N, 2007-2014



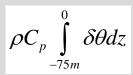
Average heat imbalance (2007-2014)

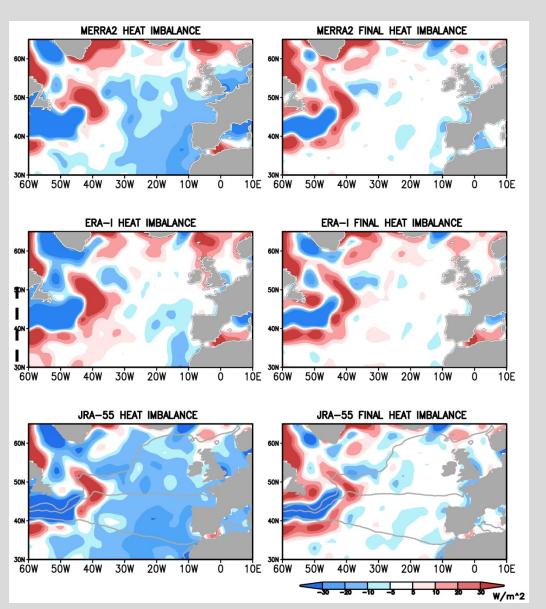
-75m



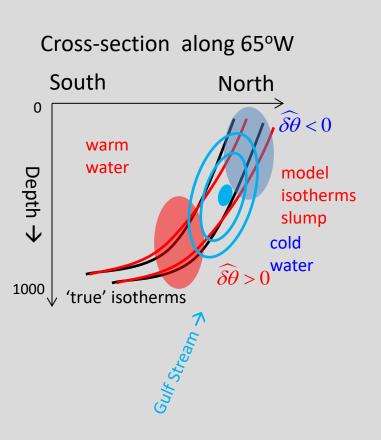
Model error dominates in western boundary currents

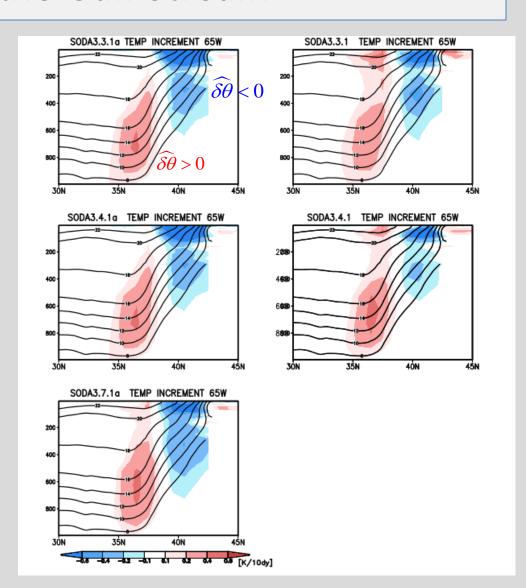
Detail from the North Atlantic





Cross-section of temperature imbalance $\widehat{\delta\theta}$ across the Gulf Stream



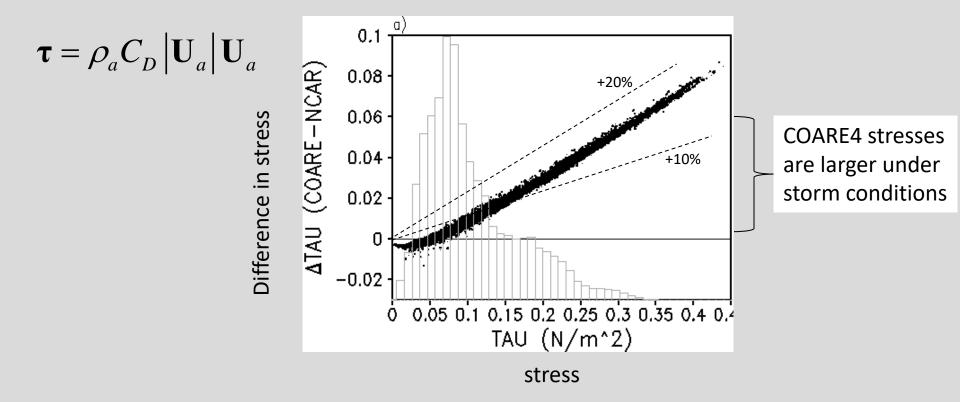


Impact of turbulent flux bulk formula: COARE4 vs NCAR

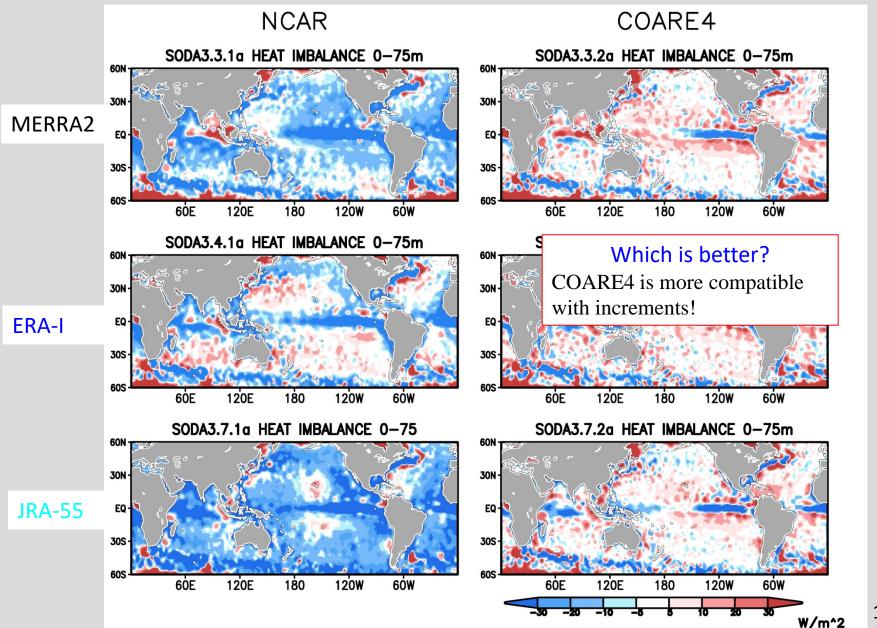
$$Q_{l} = C_{l}L\rho_{a}\mathbf{U}_{a}(q_{a} - q_{s})$$

$$Q_{s} = C_{s}C_{p}\rho_{a}\mathbf{U}_{a}(T_{a} - SST)$$

Impact of stress bulk formula *NCAR* vs COARE4

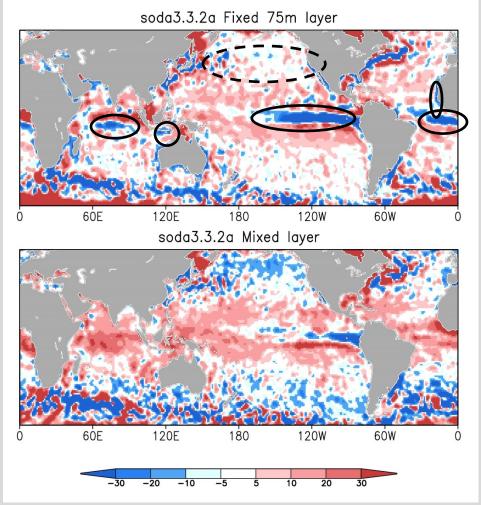


Impact of different flux parameterizations



Impact of a more realistic estimate of vertical scale of heat penetration

Fixed 75m layer



Comments

We've applied sequential DA to tackle a problem previously left to inverse methods – using ocean obs. to correct surface fluxes. It works great: it's cheap and iterative.

In playing with this approach we find:

- Most surface net heat flux misfit (error) is stored locally in the mixed layer – and it is mainly seasonal.
- Approach corrects seasonal heat fluxes to within ±5Wm⁻² except in regions of strong currents.

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- Approach also provides a way to evaluate
 - model error
 - Bulk flux parameterizations
 - Sea ice thickness